Unpublished Draft Manuscript for the

Lincoln Point 7.5' Quadrangle Utah County, Utah

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by.

Date: ~ early 1999?

(U65 Contract Mapping Project)

#### INTRODUCTION

Lincoln Point quadrangle is in the southern part of Utah County. The northwestern part of the map includes a portion of the Pelican Hills and about a square mile of the eastern Lake Mountains. Several square miles of Utah Lake, including tiny Bird Island, make up the central "area" the southern part of the map area includes Lincoln Point, the northern portion of West Mountain, Utah Lake on the west and some flat range and farm lands to the east.

Access to the northwestern part of the map area is by State Highway (7), a hard surface road from the south at Elberta, or from the north via Lehi and/or Salt lake City. This runs near the lakeshore. The Lincoln Point area is accessible via Highway, -also hard surface and which parallels the eastern part of Lincoln Point, and passes around the point and thence southward to Genola. Several dirt roads lead into the mountainous areas of the quadrangle.

Elevations in the area rise to about 5430 feet in the Lake Mountain and to 5032 feet in the West Mountain portion. The mountainous terrain, consists of gentle slopes throughout. Mountainous terrain, consists of gentle slopes throughout. Mountainous terrain, consists of gentle slopes throughout. Mountain of gentle slopes near the hills and to grazing and farming in the flat lands. Vegetation in the mountain ous areas consists mainly of grass and sagebrush. Juniper trees grave, along the highest ridges in the Lake Mountains part. On list Bird ting Island at least three varieties of deciduous trees are present. The surface elevation of Utah Lake varies depending on water usage, but is about 4490 feet (1369 meters ) in a normal year. Several dirt roads lead into the mountainous areas of the quadrangle.

The map area lies between 111 45'00" and 111 52'30" west longitude and 40 07'30" - 40 15'00" north latitude. The northwest land areas make up parts of secs 31, 32 of T6S, R1E, sections 5, 6, 7 and 18 T7S, R1E, and the southern land mass is included in parts of sections 2, 3, 9, and 16, T8S, R1E and parts of sections 6, 7 and 18 of T8S, R2E.

The climate is that of the eastern valley areas of central Utah with snow in the winter, a fairly moist spring time, and a moderately warm summer and fall with occasional rainstorms.

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Samples were collected for geochemical analyses both from the land areas and Utah Lake." Spamples from the lake, numbering 13, were obtained from a boat by driving a plastic pipe into the lake bottom sediments about one foot (0.3m) and retrieving the pipe and included samples. An east-west profile, two miles north of Bird Island, was sampled across the lake on a half mile interval. well as two samples from Bird Island, one of the beach gravel and another of a partly altered tufa outcrop at water level.

about Sample locations and others collected to the map area are

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a. Index Map of Utah Lake and Lincoln Point Quadrangle.

b. %Ca. x=25.1%; S.D.=4.9%; Rge: 8.5-31.1%; -1S.D.

c. %Mg. x=1.34%, S.D.=0.16%, Rge: 0.73-1.50%.

d. %Fe. x=0.63%, S.D.=0.07%, Rge: 0.24-1.15%.

Figure \_\_\_\_\_ a, b, c, d. a. Index map showing depth of lake in feet, and location of lincoln Point Quadrangle; b. %Ca content, upper bottom lake sediments; c. %Mg content of upper bottom lake sediments; d. %Fe content, upper bottom lake sediments. Stipple pattern is above mean content for element. ( Contours based on Sonerholm data, 1974 ).

a. %Al content: x=1.34%,S.D.=0.17%, Rge: 0.81-2.75%.

b. %Si content: x=11.83, S.D.=4.95%, Rge: 7.63-33.76%.

c. %K content: x=0.42%, S.D.=0.08%, Rge: 0.20-1.58%.

d. %Na content: x=0.10%, S.D.=0.04%, Rge: 0.06-0.33%.

Figure \_\_\_\_\_ percent content of Al, Si, K and Na in upper bottom lake sedements of Utah Lake, Lincoln Point quadrangle. Stippled pattern includes areas contaning greater than x content of element, diagonal lines: land area.

4 4 4 6 6 6 8 8 8 10 10 10 11 11 11 11.5 11.5 12 12 N N N 0 1 2 3 4 5 Miles Miles 0 1 2 3 4 5

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shown in Figure !\_\_\_\_\_

The samples were analyzed geochemically by Acme Analytical Laboratories, Ltd. of Vancouver, B.C., and in the Brigham Young University Laboratories by XRF and XRD by Dave Tingey and  $\mathfrak{P}$ . Dana Griffen, respectively.

Field work commenced in spring, 1995 and carried part time through 1995. Land areas were mapped by Proctor and Wang, and Utah Lake bottom samples planned and collected by Brimhall, Proctor and Wang.

# Previous Work

The first known written record of the area is that of Padre Escalante and Padre Dominiguez, the Spanish priest explorers in about 1776 (Auebach, 1934). The padres entered the southeastern part of Utah Valley on 23 September 1776, and reached the shores of now named Utah Lake . They named it "Laguna de nuestra Senora de la mercedide Timpanogotypes" and wrote that the lake had an outlet to the north. A member of the party, Don Bernardode de Mieray Pacheco, made a map and showed mountains at the southeast end of the lake in a position of the present West Mountain.

The John C. Fremont party camped near the south shore of the lake (May, 1844). Earlier, other U.S. Army Engineers parties surveyed for possible railroad routes. The first geologic reference to the West Mountain area is an Atlas sheet of the Wheeler Survey, 1875, on which strata of West Mountain are shown as Carboniferous and the name Wasatch applied. Geologists of the Fortieth Parallel Survey (King, Vol. 1, 1878, p. 155-248) gave the name Weber to most of the Carboniferous strata of the Wasatch Mountains but Lake Mountains was not included. In a map accompanying the report, King showed the outline of the south half of Lake Bonneville from Gilbert's map of the Wheeler Survey and the north half directly from the map made by Gilbert for King.

Other publications have appeared on this part of Utah Valley since the work of Gilbert, Butler's reconnaissance geologic map of Utah (1920, pl. IV) shows northern West Moutains and Lake Mountains as all Carboniferous stratigraphy. Davis (1925) discussed features of the Basin and Range geology and commented on West Mountain structures, but did not mention stratigraphy. Some details of the geology of West Mountain are reported by Eaton (1929, p. 71-79), who describes a simple anticlinal fold structure developed in Missippian limestones bounded by block faults on the east and west. Eardley's geologic map (1933, map 11) regrettably ended 2 miles south of West Mountain.

Bullock'5 (1949) bomple general stratigraphy and structure of Mountains and describes the general stratigraphy and structure of the mountain range. Bissell (1963) reported on the geology of southern Utah Valley, and commented on the Paleozoic rock masses in this part of the county, but did not subdivide the rock units. He

did describe in some detail the surficial units in the area. Lovering and his associates (1951) better defined and modified stratigraphic nomenclature in the Tintic mining district to the west, while Tooker and Roberts (1970) attempted a subidivision of the very thick Oquirrh Formation, earlier described by Gilluly (1932). This rock mass makes up a good part of West Mountain and also of the Lake Mountains to the northwest. In the Wasatch Mountains the east Baker (1947) described to the stratigraphy near Provo, Utah. More recently Brimhall (1973), Sonerholm (1974), Merrett (1980), completed detailed studies of the sediments and waters of Utah Lake. Machette (1989) mapped and compiled the surficial geology of the Wasatch Fault zone, which included a part of the quadrangle area southeast of Lincoln Point, as well as data on lake cycles and neotectonics of the younger sediments in the area (1988). About the same time (Currey et al, 1983) published a regional map showing the major levels of Great Salt Lake and Lake Bonneville including the area of the quadrangle. Price and Conroy (1980) reported on their study of the ground water resources of the southern Wasatch fault area of Utah. Miller (1982) mapped the surficial geology of the area, including Lincoln Point and modified earlier stratigraphic nomenclature. Two maps appeared in 1982 (Price and Jensen) describing the surface water resources of the southern Wasatch Fault, Utah. One showed location of stream gaging stations in Utah County and the range of dissolved solid concentrations in runoff at the indicated sites. An unnumbered site is shown on Benjamin slough near the highway to Lincoln Point. Witkin and Weiss (1991) mapped the most southern part of West Mountain, south of Lincoln Point quadrangle.

Unpublished works directly related to West Mountain are those of Nolan (1950), Swanson (1952), and White (1953). These describe the stratigraphy of the Oquirrh Formation, and details of the southern part of West Mountain and the general geology of West Mountain, respectively. Okerlund (1951) describes the calcite vein deposits and general stratigraphy just north of the map area in the Lake Mountains.

# Stratigraphy

The restricted bedrock areas of the quadrangle limit the exposures of the Paleozoic rocks, the oldest exposed in the area. Only portions of known Paleozoic formations crop out, except for the Humbug Formation which is fully exposed. The Oldest Paleozoic formation exposed is the Mississippian Deseret Limestone. This is overlain by the Humbug Formation, and Great Blue Limestone in the Remarkant northwest part of the quadrangle. A portion of the thick Oquirrh Group is exposed just south of Lincoln Point. South of the Mquadrangle boundary this rock unit is almost completely exposed. The older rocks have been folded, faulted, and eroded and overlain by Tertiary North Horn conglomerates and younger limestone of possible Flagstaff age. All of the indurated rocks are partly covered by surficial deposits of Lake Boneville and younger alluvial fan deposits. Utah Lake sediments and tufa cover the lower elevations of The tuba is related to thermal springs, mading near Lincoln Point and on Bird Sand

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exposed near Lincoln Point and on Bird Island. Relatively recent fine-grained calcareous sediments floor Utah Lake.

# Paleozic Rocks

#### Mississipian System

#### Deseret Limestone

Only a partial section of the Deseret Limestone is exposed within the map area. This section crops out in sections 6 and 31, T7S, R1E in the northwestern part of the map, and also in sec 7, just to the south. The beds trend similar to the overliving Humbug Formation in the northern outcrop area, but north northeasterly in sec 7 they dip moderately to steeply east. They latter show a high angle of discordance to the cherty section of the Great Blue

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Formation, and are probably separated from them by a thrust fault. The percept cherty limestone beds continue to the Utah Lake shoreline in sec 6. They are covered on the north by Lake Bonneville sediments in sec 7.

The lithology of the formation is typically gray-blue cherty limestones, generally thin-bedded. Included chert masses are brownish weatheing, nodular to layered, and as thick as six inches (\_\_\_cm). Strike length is several feet. The chert masses typically weather brown and show slight effervescence with HCl acid. As much a $\int_{-}^{-}$  feet (\_\_\_m) of cherty limestone are exposed in sec 7. To the north the exposed part is much thinner and is cut off on the east by a steeply dipping normal fault just north of the quadrangle boun#dary.

### Humbug Formation

A complete stratigraphic section of the Humbug Formation is exposed in the northwestern part of the quadrangle on the east flank of the Lake Mountains. This section consists of sandstones, orthoquartzites, interlayered limestones and cherty dolomitic beds near the base. The lowermost beds lie directly on the Deseret Limestone, and the uppermost beds are covered by the Great Blue ledge limestones.

The Humbug Formation is about \_\_\_\_\_feet ( m) in thickness as measued in section 6, T1S, R1E. The beds strike generally northwesterly and dip gently southward with outcrop development fair to good. They form a conspicous east slope of the mountains in this area.

A measured section of the formation is shown in Figure\_\_\_below.

353	C39	U12	-0.01381	0.00776	0.343
354	C40	X/A	-0.02637	0.00128	0.150
355	C40	Y/B	0.36504	0.00059	-0.730
356	C40	Z/C	0.13961	0.00043	-0.501
357	C40	U11	0.07636	0.01082	-0.493
358	C40	U22	0.06726	0.00985	0.171
359	C40	U33	0.06901	0.00935	0.343
360	C40	U23	0.01933	0.00831	-0.257
361	C40	U13	-0.02357	0.00802	0.332
362	C40	1112	-0 00194	0 00896	0 027
363	C41	¥/A	0 07726	0.00126	-0 228
364	C41	Y/R	0 39853	0 00053	0 368
365	C41	7/0	0.1355/	0.000/1	0.177
744	C/ 1	2/0	0.07204	0.00051	0.17/
367	C41	1122	0.07200	0.00931	0.134
740	0/1	1177	0.05087	0.00071	0.090
300	0/1	033	0.00079	0.00921	-0.317
309	C41	025	0.01408	0.00704	-1.094
570	641	013	0.00634	0.00812	0.655
5/1	C41	012	0.00133	0.00794	0.779
372	C42	X/A	0.18028	0.00107	-0.225
373	C42	Y/B	0.38538	0.00059	-0.573
374	C42	Z/C	0.15959	0.00042	-0.545
375	C42	U11	0.03412	0.00810	-0.556
376	C42	U22	0.07825	0.01027	0.167
377	C42	U33	0.07132	0.00919	-0.157
378	C42	U23	0.01986	0.00853	-0.373
379	C42	U13	0.01006	0.00731	0.011
380	C42	U12	-0.01055	0.00736	-0.371
381	C43	ROTX	0.02546	0.00650	3.915
382	C43	ROTY	-0.00628	0.00668	-0.941
383	C43	ROTZ	-0.00481	0.00606	-0.793
384	C43	X/A	0.65657	0.00079	-1.658
385	C43	Y/B	0.12396	0.00045	1.296
386	C43	7/0	0.20361	0.00037	-0.871
387	C43	u11	0.07168	0.00536	1.034
388	C44	111	0.07458	0.00523	1.919
380	C/5	1111	0 07162	0 00495	2 582
300	C/.6	1111	0 06500	0 00495	1 430
370	c40	111	0.0503/	0.00457	1 7/0
371	C47	1111	0.05/54	0.00431	0 253
202	C40	V/A	0.04431	0.00421	1 755
201	649	X/A	0.01199	0.00191	-1.335
394	649	1/8	0.19636	0.00084	-0.050
395	649	2/0	0.24483	0.00065	0.559
396	C49	U11	0.06854	0.00563	-0.705
397	050	X/A	0.58834	0.00104	-1.580
398	050	Y/B	0.22247	0.00049	-4.234
399	050	Z/C	0.25480	0.00038	5.929
400	050	U11	0.06236	0.00325	0.643
401	051	X/A	0.40340	0.00131	0.139
402	051	Y/B	0.20850	0.00056	0.296
403	051	Z/C	0.25134	0.00046	0.853
404	051	U11	0.09197	0.00437	-0.291
405	050	U11	0.10217	0.01120	-0.670
406	C49	U11	0.09864	0.02068	-1.285
407	C48	U11	0.05824	0.01112	0.339
408	C45	U11	0.06143	0.01046	1.092
409	C47	u11	0.08523	0.01070	3.291
410	C44	u11	0.09352	0.01059	4,111
411	C43	U11	0.10538	0.01062	5,213
412	051	111	0.10146	0.00713	7.214
416	0.1	011	0.10140	0.00113	1.214

MEAN AND MAXIMUM SHIFT/ESD: 0.644 7.214

CORRELATION MATRIX ELEMENTS GREATER THAN 0.5

382. 386. 0.6731 383. 384. 0.5335 383. 385. -0.6806

Figure Measured stratigraphic section of the Humbug Formation, Sec 6, T7S, R1E, Utah County, Utah.



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# Great Blue Formation

In the northwestern part of the quadrangle

Wo separate masses of Great Blue Formation crop out in section 6-7 and 6-31, T7S, and T6S, RlE along the east flank of the Lake Mountains. The northernmost section strikes northward and consists of prominent ledges of blue-gray massive to thick-bedded limestones. The beds dip gently westward out of the quadrangle. The southern outcrop area exposes an upper section of cherty limestones of northwesterly strike and moderately to steep southwesterly dip.

Off the quadrangle border to the west the lower ledge maker massive and thick-bedded limestones are covered by the cherty blue-gray limestone, and an interlayered section of thin beds of quartzite, shale and limestone beds of steep southwesterly dip. This unit probably represents a part of the Long Trail Member mapped to the southwest a few miles by Proctor (1954), and by

Gilluly (1932) and Bullock (1949).

The ledge limestone units represent the limestone beds currently being mined in the Larsen quarry just of the north edge of the map, which have been downdropped eastward on a normal fault.

Chip samples collected from the thick-bedded ledge limestone section were analyzed for trace metal and calcium content. Results are shown in Table \_\_\_\_\_below.

Table \_\_\_\_ Geochemical analyses of ledge maker beds of the lower Great Blue Formation, Sec. T R, Utah County.

Stratigraphic characteristics of the exposed Great Blue Formation are shown in the measured section below (Figure )

Figure. Measured stratigraphic section of the lower Great Blue Formation, Lincoln Point quadrangle, Sec T R, Utah County, Utah

TO: All Math Graduate Studednts FROM: Lonette (378-2062) 14 March 1996 DATE: RE: Spring/Summer tuition After reviewing our financial situation, it has been determined that we have extra money this year and hope to be able to offer tuition support to interested persons. Please list the mathematics courses you would like to take during these terms and return the form to me by Monday, 18th March or call and let me know. Name: Social Security Number: Courses you want to take: Thanks.

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A partial section of the Oquirrh Group is exposed in the southern part of the quadrangle, Wwhere less than a square mile of Oquirrh outcrop is present in sections 15 and 16, T7S, RIE. These outcrop masses represent the northernmost exposures of the well-developed Oquirrh expositions southward in the West Mountain quadangle. The Oquirrh beds are moderately to steeply inclined eastward and trend northward.

Common lithology in the westernmost outcrops south of Lincoln Point is minor orthoquaertzite, then a northward bearing gently inclined section of blue-gray cherty limestone of platy characteristic, These limestones are locally fossilferous, containing brachipods and some fenestellid types of bryozoa. Eastward Lake Bonneville sediments cover the bedrock.

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Hastward of the cherty limestones several hundred feet, are distinctive gray white to buff quartzite beds which strike northward and are overturned to the east. These range to <u>Goo</u> feet ( m) thick. The lower quartzite beds show good cross-bedding, as well as local convoluted beds of primary origin. The uppermost beds of the quartzite section are more regularly bedded, locally laminated, and expose local fine-cross bedding. Because of the distinctive lithology this quartzite section is mapped as avmember within the Oquirrh group.

The uppermost beds of the Oquirrh Group strike northward and are strongly overturned to the east. Common lithology is minor interbedded limestone, and more common sandstone, orthoquartzite, calcareous sandstones and arenaceous limestones. The uppermost beds are covered by Lake Bonneville sediments. Total exposed thickness of the Oquirrh Group approximates \_\_\_\_\_feet (\_\_\_\_m). Just off the map to the south and also northwest of the mapped area, a very thick section of the Oquirrh Group is exposed. *A measured* 

pection of the capped Ognitia group is included in Figure .

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Fig measure a section of exposed orguerth groups Servicion Pour and attack Guily, 10400

# David G. Wright, 12:02 PM 3/5/96 ... Fric Swenson

X-Sender: wright@hemblin.path.byu.edu Mime-Version 6 Date: Pue, 5 Mar 1996 12:02:55 -0700 To: fielding@hamblin.math.byu,edu From: "David G. Wright" Awsightemath.byu.edu> Subject: Aric Swenson Status: Jill, please make a letter for Eric Swensen for R. Weiss, Hiring committee Chair Tufts University Department of Mathematics Medford, MA 02155 Thank

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# North Horn Formation

A distinctive conglomeratic unit crops out on Lincoln Point ridge near the south boundary of the qudrangle. We identify this as probable possible North Horn Formation. It consists of a single bold outcrop - more than 25 feet thick of tannish to pinkish weathering boulder conglomerate and pebbly sandstone. Clasts range to 6 inches ( cm) or more in diameter and are mainly subrounded quartzite boulders with accompanying gray-blue limestone boulders and cherty limestone clasts. The mass crops out in sec\_\_\_\_, T7S, R1E. In this same area, and to the north toward Lincoln Point, the surface is partly covered with large bounders of quartzite and limestone as much as 6 feet (2 m) and more in diameter. These appear to be reworked clasts from the underlying North Horn Formation. They are randomly distribted over the surface but are abundant and conspicuous.

Small bands of blue-gray limestone clasts and some quartzite clasts, cemented by calcite and/or reddish calcite, crop out in the same area. These are visible on the east side of Lincoln Point ridge, on top of the redge and a lesser number of outcrops are exposed on the western side toward the top of the ridge in sec\_\_\_\_, T7S, R1E . Included clasts are up to six inches ( cm) or more in diameter, are are cemented by reddish to gray calcite. These lowlying outcrops probably repesent bedded layers of the North Horn Formation. Most of the underlying sediments, of similar composition and texture, is covered by slope wash and reworked Lake are Bonneville sediments. Is goilly relied cartward

Thickness of the North Horn Formation is not directly available, but the scattered outcrops of the flat-lying formation, and residual boulders derived from it suggest a thickness of as much as 500 feet on the northern part of West Mountain near Lincoln Point.

#### Quaternary System

Surficial deposits of the Lincoln Point quadrangle were deposited within the time-event framework so well described by Machette (1989). Several mappable units are identified and mapped by him, including Pre-Bonneville alluvial fan complexes, reworked by Lake Bonneville, Pre-Bonneville slope wash also reworked by Bonneville waters, Lake Bonneville gravels, post-Provo level alluvial fan sediments, Utah Lake clays and silts, tufa deposits, and the youngest sediments in Utah Lake, and alluvial deposits in small washes and in stream beds. Most of these surficial units are identified within the map area and mapped on the larger map scale.

The age relationships and character of the deposits within the map area are shown in Figure \_\_\_\_ below. A description of each of the mapped units follows.

TO: All Math Graduate Students
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RE: Spring/Summer tuition

After reviewing our financial situation, it has been determined that we have extra money this year and hope to be able to offer tuition support to interested persons. Please list the mathematics courses you would like to take during these terms and return the form to me by Monday, 18th March or call and let me know.

Name:
Name.
Social Security Number:
Courses you want to take:
Thanks.
Catharan

An older series of boulders, pebbles, sand, silt and soil occurs on both northern West Mountain and on the east flank of the Lake Mountain. On *Mace* north West Mountain these consist of partly reworked erosional materials of the Paleozoic bedrock, *case* Mear Lincoln Point of the Tertiary-Cretaceous North Horn Conglomerate. At this latter location large residual boulders to 6 feet (2 m) or more in diameter of quartzite, limestone and cherty limestone occur randomly distributed along the slope. These lie upon the bedrock conglomerate. Smaller size boulders, pebbles, and some sand and silt also occur in conjunction with the larger boulders. Both the larger diameter boulders and smaller diameter materials have been partly reworked by Lake Bonneville waters both above and below the prominent Provo Level terrace of the ancient lake. Former lake levels are faintly to prominently displayed in this sedimentary cover material.

Boulders of bedrock composition, and up to 6 inches in diameter, as well as pebbles, sand and silt debris occur on the east slope of the Lake Mountains in the northwest part of the map area. These sediments are derived from the underlying Paleozoic bedrock. All have been somewhat reworked by the waters of Lake Boneville both above and below the prominent Provo Level terrace of the former lake. These are mapped as re-worked Pre-Bonneville slope wash deposits.

# Pre-Bonneville Alluvial Fan Deposits

Prior to the deposition of the transgressive Lake Bonneville sediments, a series of relatively large alluvial fan complexes developed along the flanks of both the Lake Mountains and northern West Mountain. These consist of limestone, cherty limestone, orthoquartzite, and sandstone bould of pebbles, sand and silt. They show a convex form outward from the mountain front, and are readily recognized by the geometric form of the contours. Two such older alluvial fan deposits occur along the western slope of north West Mountain in secs. 16 and 9, and in part of sec 10, T7S, RIE. Both masses are expressed top ographically by the contours which are concave to the bedrock surface and range from above 4800 feet ( m) to the present level of Utah Lake at 4490 feet (\_\_\_m).

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Composition of these surficial sediments is that of the nearby bedrock whose derivative it is. While thickness of the fan material is not directly observable, the sediments have an elevation difference of at least 300 feet, and if lens-shaped may be as much as \_\_\_\_feet thick.

# Lake Bonneville Gravels, Sands and Silts

Transgressive water of the rising Lake Bonneville locally reworked colluvial materials and erosional debris from the mountains slopes. Gravels were small cobbles, pebbles and sand were carried by and redeposit by currents at the upper levels of the lake at approximately 5125 feet ( m) and below. Best exposures of thick Lake Bonneville gravel deposits are in the



STECHLY TILTEN (FINDLO) LAKE BONNEWILL ORADU IS, LITTLE COVE, Northwood by quadrangle. Sec TK

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Little Cove area, Lake Mountains in the northwest part of the map area. Here up to 300 feet or more thickness of poorly layered gravels accumulated. They are restricted to the Little Cove area and clearly lie directly on the Paleozoic bedrock. Individual clasts are subrounded to discoidal and are supported by cobbles, pebbles and sand in a matrix of sand and silt. Some cementation by carbonate deposition is present. Thin beds of sands, moderately sorted are interlayered. The grain size of the overall deposits diminishes to the east.

#### Utah Lake Bottom Sediments

Numerous studies have been made on the bottom sediments of Utah Lake, including the part within the Lincoln Point guadrangle (Bissell, 1942, 1963; Bradshaw et al, 1969; Brimhall, 1973, 1976; Sonnerholm, 1974; Bingham, 1975; Bushman, 1980; Sundruck et al, 1970; Brimhall and Merritt, 1981; Hobson, 1994). Bissell (1963, p. 122) writes that a "clay unit almost completely borders the present 15 edge of Utah lake and also underlies most of the present lake The clay is dark gray, dark blue gray, and locally black botttom. owing to orghic matter....". He further records that: "The upper 5 to 8 feet of sediments on the bottom of Utah Lake is light-tomedium gray silty clay that is fairly well sorted, poorly compacted, and contains 50 to 60 percent water". This, as he notes, grades downward into the more compact medium to dark-gray clay which extends to a depth of 15 feet. Below this, he describes a 10 foot thick bed of well-sorted, dark gray to black compact silty clay with 33 to 43 percent water content, which becomes sandy at a depth of 25 feet. In addition he (op cit, p.122-3) describes a tufa unit near Lincoln Beach and along the west side of north West Mountain along the lake shore. It occurs slightly below the high-level of the lake and extends outward from the shore. The tufa also makes up the base of Bird Island, and probably occurs as a shield around it. Hobson (1994) also describes the same rock type and outlined it by sonic surveys in this part of Utah Lake. Cottam earier (1926) suggested that the tufa of Bird Island is deposited on a pedestal of quartzite.

The current geologic mapping confirms the tufa of Lincoln Beach, and that westward and southward along the shoreline. Beach gravels are also partly cemented by it. The contact mapped by Bissell (pl. ) is used because he observed the deposit at a lower lake level than was available during the present field work.

A sediment sample profile was run across Utah lake two and a half miles north of Bird Island. Twelve samples were collected of the uppermost light gray clay-silt unit of the lake bottom. Individual samples were about 2000 to 2500 feet apart and were located on the map by utilizing the GPS system in a small boat. Samples were obtained by driving PVC pipe, one-inch in diameter into the sediments, then recovering the pipe and included sediment. The end of the pipe was sawed off with its contained sediment sample, and capped at both ends to preserve both water and claysilt for analysis. Additional samples were collected about two miles WNW of Bird Island and from the bedrock tufa of the island, and from **WAP** beach gravel.

The twelve samples were analyzed by XRF in the Brigham Young University laboratory by analyst David Tingey. Results are shown in Table \_\_\_\_\_. Several of the collected samples were also analyzed for trace metal content by Acme Laboratories, Ltd., of Vancouver, B.C., a certified commercial laboratory. Analytical

Bissel 1963

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results are shown in Table \_\_\_\_\_.

Distribution Patterns of Major Elements in Utah Lake Sediments

Earlier, Sonnerholm (1973) carried out an extensive sampling program of Utah Lake sediments. He collected 139 samples on an approximate 1 mile (1.6km) grid. These uppermost bottom lake sediment samples were analyzed by atomic absorption for Ca, Mg, Fe, Al, Si, K and Na. Results were plotted and Trend Surface maps were prepared for each of the elemental results.

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Chemical results from samples within the quadrangle boundaries were statistically reworked for the mean (x), range, and standard deviation from the mean (Table\_\_\_). Using these statistical values we prepared contour maps, based on one standard deviation to two standard deviations from the mean \_\_\_\_\_\_ threshold values, from greater than two standard deviations from the mean as anomalous contents. Individual maps for each of there seven metal content plots are included in Figures \_\_\_\_\_. The interpreted threshold and anomalous metal content areas are shown by pattern on each of the maps. Individual samples considered threshold or anomalous in metal content are also shown by symbol on the maps.

As a generalization, the mean contents of Ca and Mg lie mainly in the central part of the water covered area within the quadrangle. Al and Fe show a distinctive SW and south bearing pattern of lower than mean content in the central map area. The Si, K and Na patterns resemble each other, with the above mean contents trending generally north and slightly eastward. All low metal contents occur in the central part of the map area. The marginal high content areas tend to parallel the Utah Lake shorelines.

Sonnerholm (1973), using the analytical chemical results of the major elements, developed nine normative minerals including calcite, magnesite, gypsum, quartz, clinochlore, illite, hematite, halite and sylvite. He concludes (1974) that normative carbonates (calcite and magnesite) are the most abundant normative mineral types, the two combined representing 90% or more of the samples collected. We concur in his assessment.

Our twelve profile samples collected on an east-west traverse two miles north of Bird Island, are shown in the profile of Figure\_\_\_\_\_. This compares the analytical results for Ca, Si, Al and Fe. The profile clearly shows the higher CaO content (most likely representing calcite content) in the central area, and less abundant SiO2, mainly quartz, in the same general position as the calcite. Contents of SiO2 increase toward the west and east, or toward the shoreland areas. This correlates well with the work of Sonerholm. Alumina, probably represents clay and silt fractions of the sediments, and shows a fairly even profile, but with a slight drop in content toward the shoreland areas.

The profile set of samples also includes geochemical analyses

for several trace metals and transition elements, S, F, Cl, Zn, Sr, Cu, Pb, Ni, As and Cd, analyzed in parts per million (ppm) content. Thus are plotted in Figure \_\_\_\_. Distinctive patterns are shown for S, Zn, Fe, F, and Cl contents. Each has a sharp rise toward the east. The Cd content shows a high content near the center of the profile then an abrupt rise in content among the easternmost samples. Several factors might explain these results. For example, a source to the east and north, i.e., Geneva Steel plant, or the waters draining into the lake from the east with their special metal contents such as American Fork stream, Provo River, Hobble Creek and Spanish Fork Creeks. Another factor might be metal additions from the thermal spring areas present in the lake, but the instance

The gologic profile which accompanies the geochemical profiles shows the very shallow character of the lake, the general location of known faults in the lake sediments, and the thermal spring areas. The content of Zn, F, Cl and possibly S in the area of the thermal springs, begins to increase in content at least one half mile west of these plotted features and continues to increase at least one half mile east of them. There is a suggestion that the thermal zones are related to the increase in content of these elements.

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Two samples of tufa were collected from Bird Island. A third sample of tufa was obtained from the tufa outcrop on Lincoln Beach. In both areas the tufa is concluded to have been deposited from nearly thermal spring activity. The geochemical analyses show a lower content of heavy metals (\_\_\_\_\_vs\_\_\_). One of the Bird Island samples showed limonitic staining, possibly the result of oxidation of pyrite in the rock. The As content is about 13 times higher (189 vs 14), than the mean for the other fourteen collected samples. Other trace metals Cu, Pb, Zn, Ni, Fe, V, Cr, Ba, B are in lesser quantities than the mean for the uppermost lake sediment of the profile. Indeed, the trace metal contents are quite similar to those known (Proctor and Wang\*, unpublished data) for the large calcite veins currently being mined just north of the northwest map boundary.The the are believed to be of thermal water origin..

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					Table	C					
		>	KRF Major	Element /	Analyses of	f Utah lake	Bottom S	amples,			
		F	Profets	,	Lincoln Po	int Quadra	ngle, Utah	Co., Utah			
0 I N	0.00	41000	F-000	0.0	14-0	N-00	KOO		DOOL	TOO	
Sample N	SI02	AI2O3	Fe2O3	CaO	NIGO	Na2O	K20	MnO	P205	1102	LOI
UL-141	25.09	5.37	2.42	29.00	4.06	1,05	0.92	0.03	0.02	0.33	30.19
UL-143	22.66	4.92	2.18	30.95	4.07	1.03	0.84	0.03	0	0.31	31.14
UL-144	22.86	4.88	2.16	30.77	4.03	1.01	0.85	0.03	0	0.31	31.14
UL-145	23.06	4,89	2.16	30.99	4.03	0,97	0.84	0,03	0	0.30	31.31
UL-146	22.65	4.82	2.12	31.15	4.01	0.96	0.83	0.03	0	0.30	31.09
UL-148	21.21	4.67	2.10	32.05	4.12	0.88	0.79	0.03	0	0.28	32.22
UL-150	20.37	4.56	2.13	32.66	4.11	0.92	0.76	0.03	0	0.28	32.27
UL-153	34.21	5.31	2.08	24.72	3.33	1.22	1.05	0.03	0.04	0.34	25.61
UL-154	21.95	4.80	2.15	31.39	4.02	1.01	0.81	0.03	0	0.29	31.26

# Selected Minor Element XRF Analyses and of the Selected Minor Element XRF Analyses and of the Selected Minor Element XRF Analyses and the Selected Minor Element Analyses and the Selected Minor Element XRF Analyses and the Selected Minor Element XRF Analyses and the Selected Minor Element Analyses and the Selected Minor Element XRF Analyses and the Selected Minor Element Analyses and the Selected

Sample N	F	S	CI	Cu	Pb	Zn	Ni	Sr	Ba	As	Ca	Zr
UL-141	397	3626	398	24	22	130	17	1089	271	13	0.9	60
UL-143	316	3176	306	25	20	109	13	1186	258	15	0.6	52
UL-144	289	3004	256	21	19	103	14	1185	260	13	0.8	53
UL-145	265	2777	250	24	20	102	13	1193	254	12	0.6	53
UL-146	260	3267	255	19	18	93	13	1188	256	16	0.5	53
UL-148	208	2751	252	19	18	88	13	1246	253	17	1	49
UL-150	194	3092	289	19	17	84	12	1258	248	14	0.2	48
UL-153	314	2789	286	27	18	87	14	971	296	11	0.4	98
UL-154	263	3401	292	25	19	96	13	1191	256	14	0.4	52

\* Analyses: Department of Geology, Brigham Young Universit rsity, Pro ovo, Utah, David Tingey, analyst.

\*\*Range(ppm) of Sc(L1), V(41-48), Cr(12-14), Ga(11-13), Rb(L1-1), Y(9-11), Nb(8-12), La(6-11), Ce(26-35), Nd(21-26), Sm(7-8), Th(6-Sm(7-8), Th(6-7), and U(4-12) included here but not in table above.

\*\*\*Analys es by Acme Analytical Labs. LTD., Vanc ncouver, B.C., Canada.

Miner alagical Characteristics Unput the Loke terminis (V) were analyzed by X-R. Thereman anappus by Q Dan a Suffar & The Department of Fish gy at Bright Jon, Seminity, Results an indicated in the x-ray patterns of figure to the X-R pattoms indicate, the sample consest many of Calcul and quaity. D. Ingoys XRI annyous of mine of these samples confirmes abound and Dog and Cal entered, 12000 a high 201 response, most tidely Don. Druffen suggester more questy in UL-153, the SAF andyour enderales over 34 % 5.00, all other somples songelfun 20 % to 25% 510 = Cac and hos were hers for UL-153 [~24.7% fur the 25,6%) comprand to a sample of (~28% -32,6% Cal and 30, 15% - 32.27 %), specific specific under over i condified, some mine product supported day, the KRF analysis confirm Alsog with a range of 4.8-5,37 % and Digit cout in F. Og (2.1-2. 43%) with S from 2777- 3626), possibly alled to fire granied synte. She substitution & Sp in the lacente slunder is privible because of the pumer of 54 in each of the samples 1 ansing from 951 amoun to as un a ar 1250 por.

And Naso (0.88-1.22%) and K20 (0.76-1.05%) are selal and montants and J The Dampdos and year and matery sensitions about and Paul-

Very sorry to have taken so long to get these done, but here they are. As you can see, the sediments are virtually all calcite and quartz. Although getting quantitative about it would require substantially more work, I suspect that there is more calcite than quartz in all of the samples except UL153, which has more quartz. UL153 is also unusual in that there are two unidentified peaks at about 27° and 28° 2-theta, but with only two peaks to go on, I can't say what it is. The 28° peak shows up in UL149, also, but the 27° peaks does not, so the two might not even belong to the same mineral. Most of the patterns also have a small blip around 9°—that's about 10Å, so it could be illite, but nailing that down would take some clay work, and I don't know what is important to you for this project.

BYU Department of Geology XRF analyses - DAVID TINGOY ANALYST LITAH LAKE Bo Hom Sodiments- EN PRONIE North of Bird Island.

		UL-141	UL-143	UL-144	UL-145	UL-146	UL-148	UL-150	UL-153	UL-154
DATE 1996 wt. %	A	06-Mar	A 06-Mar	A Ub-Iviar	A U6-IVIAr	A Ub-Iviar	A UG-IVIAr	a ub-iviar	1 06-iviar	M Ub-Iviar
SIO2		25.09	22.66	22.86	23.06	22.65	21.21	20.37	34.21	21.95
TIO2		0.33	0.31	0.31	0.3	0.3	0.28	0.28	0.34	0.29
Al2O3		5.37	4.92	4.88	4.89	4.82	4.67	4.56	5.31	4.8
Fe2O3		2.43	2.18	2.16	2.16	2.12	2.1	2.13	2.08	2.15
MnO		0.03	0.03	0.03	0.03	0.03	0.03	0.03	0.03	0.03
MgO		4.06	4.07	4.03	4.03	4.01	4.12	4.11	3.33	4.02
CaO		29	30.95	30.77	30.99	31.15	32.05	32.6	24.72	31.39
Na2O		1.05	1.03	1.01	0.97	0.96	0.88	0.92	1.22	1:01
K2O		0.92	0.84	0.85	0.84	0.83	0.79	0.76	1.05	0.81
P2O5		0.02	0	0	0	0	0	0	0.04	0
sum		68.3	66.99	66.9	67.27	66.87	66.13	65.76	72.33	66.45
LOI		30.19	31.37	31.14	31.31	31.09	32.22	32.27	25.61	31.26
total		98.49	98.36	98.04	98.58	97.96	98.35	98.03	97.94	97.71
ppm										
F		397	316	289	265	260	208	194	314	263
S		3626	3176	3004	2777	3267	2751	3092	2786	3401
Cl		398	306	256	250	255	252	289	286	292
Sc		< 1	< 1	< 1	< 1	< 1	< 1	< 1	< 1	<1
V		48	44	43	41	42	42	41	45	43
Cr		(14)	(13)	(13)	(13)	(12)	(12)	(12)	(13)	(14)
Ni		17	13	14	13	13	13	12	14	13
Cu		26 24	26 25	2 <b>5</b> 21	25 24	2/ 19	Za 19	<i> 8</i> 19	23 27	2# 25
Zn		130	109	103	102	93	88	84	87	96
Ga		12	12	11	12	13	12	. 11	13	12
Rb		< 1	< 1	< 1	< 1	< 1	< 1	< 1	1	< 1
Sr		1089	1186	1185	1193	1188	1246	1258	971	1191
Y		11	10	10	10	10	9	9	11	9
Zr		60	52	53	53	53	49	48	98	52
Nb		10	10	10	10	11	11	9	12	8
Ba		271	258	260	254	256	253	248	296	256
La		11	8	9	6	7	11	8	8	9
Ce		33	35	30	26	34	32	30	31	38
Nd		26	23	23	23	23	22	21	22	24
Sm		7	8	7	7	7	8	7	7	7
Pb		27 22	27 20	19	20	18	18	17	18	10
Th		7	7	7	6	6	6	6	6	6
U		7	9	10	6	5	4	8	5	12

LOI - loss on ignition (CO2, H2O)

() - approximate analysis



# Rates and Character of Sedimentation of Utah Lake Bottom Sediments

Investigations on the rate and character of sedimentation in Utah Lake include those of Bissell (1942), Brimhall (1973), Bingham (1975), and Bushman (1980) and others. Brimhall (1973) studied a core sample taken from the lake bottom two miles west of the shoreline adjacent to Geneva Steel(6, This' site is north of the quadrangle. The sample came from the upper surface of the lake sediments and was 520 cm in length. He determined the variations in concentrations of Ca, Fe, P, N and Al; and suggested that the profile reflects some major historical events such as significant fluctuations of the water level and air and water pollutions of the lake. He concluded that the average sedimentation rate was 3.3 cm. per year for the past 40 years (1933-1973) and 2.6 cm per year, 1885-1935. The lower section of the core indicated that some 200-300 years ago the lake was at a very low level and had a high salinity content.

Sonerholm's study (1974), based on 139 contemporary bottomsediment samples of the lake and collected at projected section corners, involved chemical analyses of the major elements Ca, Mg, Si, Al, Fe, Na and K and the determination of normative minerals from the chemical data. The normative minerals included calcite, magnesite, quartz, clinochore, illite and hematite. Combinin g the data into four groups: total carbonate, total clay, quartz and hematite, a sixth degree trend surface map was prepared for each of the groups. **Mased for the map**, the suggested higher carbonate content in the northern and southern portions of the lake and quartz most abundant in the mid-area. Total clay and normative hematite distribution is more complex and probably reflects the influence of various streams entering the lake. He ddid not indicate rates of sedimenation.

Bingham (1975) emphasized grain-size analyses, organic carbon content, carbonate content and clay types for recent sedimentation, i.e., in the top few inches (cms) of the lake sediments. Sixtv samples were collected on east-west traverses about two miles apart, with the rows two miles apart north south. About 19 of these samples fall with Lincoln Point quadrangle. He concludes that calcium carbonate is the most abundant sediment and ranges from 6-73 percent. Clastic grains "make up essentially the remainder", and these are from rare granule size, with silt the most abundant. Coarser fractions, granule to very fine sand, are near mouths of and southerly of major streams entering the lake, and silts and clay "are more abundant in the center portions of the lake (i.e., Lincoln Point quadrangle area). Illite in the form of discrete mica is the dominant clay type, with kaolinite and mixed layer illite-montmorillonite also present. Organic material is highest in isolated bays and sloughs. Much rarer sediments near river mouths include small amounts of mica, clinkers, coal, feldspars, silicified oolites, clay chunks, and small rock fragments" His trend surface maps (his figures 4-9) clearly show the weight log percent of coarse sand through clay fractions, the trend surface

**ROCHELLE VAN ORMAN** ALYSHIA ALLEN **ALYSHIA ALLEN** BOCHELLE VAN ORMAN LESLIE CHRISTIANSEN LESLIE CHRISTIANSEN AMANDA BUTLER AMANDA BUTLER JACT DAVIS/ JACI DAVIS KRISTINA FOLAUMOELOA KRISTINA FOLAUMOELOA DARSI JOHNSON DARSI JOHNSON LIESL LEGORBURU LIESL LÈGORBURU TIFFANY PORTER **TIFFANY** FORTER JULIE RASMUSSEN JULIE RASMUSSEN KRIS/TIN/SPRINGE/R **KRISTEN SPRINGER** BRIGETTA KØENIG BRIGETTA'KOENIG ALANYA/KOENIG KIM QLARKE KIM/CLARKE ALANYA KOENIG Hannah Van Orman HANNAH VAN ORMAN MEAGAN BISHOP MEAGAN BISHOP



maps of organic material content (his figure 10) and acid soluble (carbonate) content. He cites Brimhall (1973) on rates of sedimentation. In relationship to the Lincoln Point quadrangle lake covered area, which is mainly the central lake area of his study, he concludes that the sediments contain very minor amounts of coarse clastics, usually less than one percent, clay size fractions approximate 15-21 percent, and about 10-19 percent of silt size particles. Calcium carbonate is very abundant and makes up the major part of the sediments at 61-73 percent. This is in general agreement with our own findings.

Brimhall and Merritt (1981) indicate a transition zone of loose sediments above firm lake bottom sediments which consist of about 0.5 m of "thin to thick soup". Bushman (1981) cites Fuhriman and others (1981) and points out that near Lincoln Beach, near the southern end of the lake, the waters are almost "always milky in appearance becauseof calcium carbonante being precipitated there in the form of colloidal and silt-fized crystals that remain suspended  $\nu$ for a long time". In this area beach gravels show a tufa-like cement coating, and a well-developed tufa layer is **present** around Lincoln Point. Furhiman and others (1981) calculate that calcite was precipitated at the relate of 0.2 mm per year, and that 25 percent of the total sediments and "35 percent of the calcite was deposited in the lake by mineral precipitation."

Once a clear-water lake, in the memory of living persons (Bertrand Harrison, cited 1981), and with abundant vegetation near shore, and a carpet of vegetation growth on its bottom, the lake has changed from a "clear sulfate lake:".....to become a "bicarbonate and eutrophic to dystrophic lake, with high colloidal turbidity when ice covering is lacking." By 1926 (Cottom) notes that the "lake is astonishingly free from aquatic vegetation....The disappearance of vegetation in the lake followed the introduction of the herbivorous fish-the German carp. With no enemies to impede its enormous reproductive powers the carp was soon master of the lake."

Bushman (1981) concludes that there "is good evidence that the average annual amount of detrital sediment carried into Utah Lake has increased since 1849, the time of white man's settlement of the valley. He cites deforestation and overgrazing on the eastern and western margins of Utah Valley, plus highway construction and railroad grades into Spanish Fork Canyon. Chemical changes in the lake waters have also occurred. Bissell (1942) suggests a recent increase in the amount of calcium carbonate based on its decreasing content in depth in the bottom lake sediments. Drainage canals cut in the lowlands permitted the drained water to leach salts present, decreasing salinity up to 7000 ppm (Jones, 1974) and increasing the salt content in the soils closer to the lake. Irrigation water is also cited as a major additive of salt to the lake (Turley, 1969). The latter author states that the salt content at the mouth of Spanish Fork River is three times that in the river where it is taken out for irrigation. Later Fuhriman and others (1981) suggested a range in total dissolved salts of 700 to 1000 mg/1 during average inflow years ot the lake.

After careful consideration of the abundant data available Bushman (1981) cited Brimhall's rate of sedimentation (1972) at 2.6 cm/year for 1885-1935 and 3.3 cm/year for 1935-65, and a later study of Brimhall and Merritt (1981) with a revision to 2 mm/year post-settlement rate. Varves were sought by Bolland (1974) as a Berbarry means of yielding more precise data, but none was found.

Using pollen from 45 plants to determine the "distinctive appearance and greater likelihood of being transported into the lake", Bushman (1981) used portion of a 5 m core taken in 1972 by Robert Finley Borland, a graduate student from the University of Utah, from Goshen Bay in the southern part of the lake. Slices of the core, 2 cm thick were cut every 10 cm, except the top 10 cm where 5 samples were cut. A total of 58 samples were cut, to extract pollen, with 75 thin sections and 41 processed prepared and 67 of these studied. He concludes, among other details, that the appearance of abundant algae in several samples, sugests an algal bloom in the past. Also, Taraxacum officinale, the common dandelion, was the plant by which an "approximate age level was determined from the core studied". This plant was considered to have been introduced in Utah Valley in 1849. Found in the upper 170-172 cm of the core, the rate of sedimenation from 1849-1972 is 1.38 cm per year. He states that this rate is probably the most reliable and suggests it is of the same magnitdude as reported in studies of lakes similar to Utah Lake, each of which experienced increased rates of sedimentation because of human activities in the environment. He also notes that the greatly increased calcium carbonate content in the uppermost sediments as cited. by Bissell (1942) and others needs further study.



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								1			
H22	-0.14938 0.00106	0.61366 0.00050	0.09379 0.00039	1.00000 0.00000	0.08000	-	/	/			
C23	0.02271 0.00102	0.64673 0.00052	0.08934 0.00036	1.00000 0.00000	0.04504 0.00761	0.07267	0.03551	0.00980 0.00655	0.00636 0.00639	-0.00314 0.00703	0.05107 0.00459
H23	0.04287 0.00102	0.64451	0.11976 0.00036	1.00000	0.08000	alout description of the second second					
C24	-0.24873 0.00096	0.61162 0.00049	0.01766 0.00043	1.00000	0.03784 0.00680	0.03719	0.09656	-0.01725	-0.01811	0.00410	0.05720 0.00498
C25	-0.26030 0.00130	0.61020 0.00061	-0.03282 0.00038	1.00000 0.00000	0.08895	0.08238	0.06463	0.01041 0.00780	-0.03104 0.00843	-0.00873	0.07865 0.00590
H25A	-0.33402 0.00130	0.59072 0.00061	-0.04096 0.00038	1.00000	0.10000			K AN THE SHARE SHOW AND A SHOW AND	- Andrew Market	/	
H25B	-0.19186 0.00130	0.58998 0.00061	-0.04492	1.00000 0.00000	0.10000 0.00000			- Socretaria	and the second		
H25C	-0.26180 0.00130	0.00061	-0.04379 0.00038	1.00000 0.00000	0.10000 0.00000		Non-March and Statistics of the State	excertion and the		Constant and a second second	
C26	-0.34761 0.00111	0.65215 0.00060	0.03522 0.00051	1.00000 0.00000	0.04295	0.07575	0.14349 0.01479	-0.00268 0.01027	0.00242 0.00920	-0.00357	0.08740 0.00675
H26A	-0.42690 0.00111	0.63878	0.02683 0.00051	1.00000	0.10000					and the state of the	
H26B	-0.33477 0.00111	0.69044	0.02393 0.00051	1.00000	0.10000 	Janhan og han som det som en samte som	STATISTICS OF THE PARTY OF THE	and the California of the California	Variation and the second second	*	
H26C	-0.34183 0.00111	0.65289 0.00060	0.06647 0.00051	1.00000 0.00000	0.10000 0.00000			5	2		
C27	-0.26282 0.00109	0.54945	0.03426 0.00042	1.00000 0.00000	0.03499 0.00772	0.04775 0.00870	0.09192	0:00169 0.00733	0.01961 0.00746	-0.01058 0.00733	0.05822
C28	-0.36552 0.00108	0.52677	0.05402 0.00045	1.00000	0.04364 0.00812	0.04928 0.00905	0.11085 0.01202	-0.00168 0.00824	0.00161 0.00786	0.01622	0.06792
H28	-0.43066 0,00108	0.55277 0.00055	0.06121 0.00045	1.00000 0.00000	0.08000					and a start of the	
C29	-0.37956 0.00110	0.47051	0.06616	1.00000 0.00000	0.03409 0.00784	0.05380 0.01006	0.13556 0.01273	0.03748 0.00915	0.01842 0.00791	-0.00414 0.00794	0.07448
H29	-0.45352	0.45723	0.07948 0.00046	1.00000 0.00000	0.08000		مىرمى	- Martin Martin	-		
C30	-0.28734 0.00107	0.43316	0.05906	1.00000 0.00000	0.03940	0.00912	0.06215	0.00480	0.00320	-0 <del>.</del> 00288 0.00750	0.04966 0.00491
C31	-0.18385 0.00112	0.45082 0.00054	0.03876 0.00043	1.00000	0.04304 0.00873	0.04217	0.10228 0.01064	0.00262 0.00755	0.02145 0.00797	0.00773	0.06250
H31	-0.11919 0.00112	0.42397 0.00054	0.03289 0.00043	0.00000	0.08000			/		5	
C32	-0.17102 0.00111	0.50893	0.02714	1.00000 0.00000	0.05230 0.00892	0.04679 0.00879	0.07706	0.00022	0.02110 0.00711	0.01033	0.05872 0.00524
H32	-0.09694 0.00111	0.52164	0.01359 0.00040	1.00000	0.08000 0,00000			Contraction and the second of	a fear		
033	-0.28789 0.00070	0.37431 0.00037	0.06772 0.00028	1.00000 0.00000	0.05788	0.05449	0.09552	0.01556 0.00514	-0.00013 0.00492	-0.02520 0.00493	0.06930 0.00364

#### STRUCTURES

#### Introduction

and The Lincoln Point quadrangle consists of about 12% land area of two parts: one the northwest quarter lies north and west of Pelican Point and is some 1.25 square miles in area; the second part is south of Lincoln Point in the southern part of the quadrangle and to the east. It comprises some six square miles. There is Limited exposure of bedrock in each of these areas which are separated by several miles.

The geological structures in the land areas reflect earlier results of compressive stress, most likely related to the Early Cretaceous Sevier compressional orogeny. Later faults of extensional origin are present in the bedrock, and Very young extensional features visible in the Utah Lake bottom sediments probably related/ to Basin and Range faulting of Tertiary and younger age. The structural characteristics and evolutionary relations for the three different units: Pelican Point south, Lincoln Point, and the Utah Lake unit follows.

#### Pelican Point South

Strata of the Deseret, Humbug and Great Blue Formations of the Mississippian System are exposed in the northern part of this unit. Attitudes of the strata are fairly gentle, and range from a few degrees to locally more than ten degrees. Regionally they are part of the eastern limb of the Lake Mountain syncline (Bullock, 1949). Based on the geology, the unit remesents a secondary assymetric anticline cored by Deseret Formation. The limbs consist of Humbug and Great Blue Formations. The west limb is of Great Blue and the underlying Humbug Formation. Both dip gently southwestward. The dip of the Humbug Formation gradually changes to southwesterly. The east limb is cut by a north trending, steeply east dipping longitudinal fault down to the east, and the down faulted block mainly covered by surficial materials. The anticline plunges southward and the nose is cut by a fault ( F1 ) of easterly bearing. The south block consists of strate of the Deservet, Formation and the Great Blue Formation. Strata of the Descret Formation on the east limb dip, 38 -68 southeast, While strata of the Great Blue Formation on the western limb dip southwest and at variable inclinations. The Humbug Formation, between the Deseret and the Great Blue Formations, is not exposed because of fault (over) displacement.

The fault structures of the rock unit consists of flar types:

cuts 1) A transverse EW shear fault, F1, which crosses the middle part of the fold. Evidences for the fault are not readily observed visible in the field, but we recognize the discontinuous lithizones, inconsistent bedding attitudes, and different patterns of structure on the two limbs, The shear character of the fault is inferred on

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the basis of the structural relationships among the bedding attitudes, the fault and the fold.

2) A diagonal northwest bearing fault, F2, is traceable for about 800 feet in the mapped area. It continues northwestward into the adjoining quadrangle. A major calcite vein is well-developed along this fault. On the map boundary in a quarry, the fault plane strikes N50 W and dips 70 NE. Displacement is normal and down to the north. In plan view the fault cuts the strata, and while the displacement is small, it shows apparent right lateral displacement *fabrit freque*t.

3) A thrust fault most likely exists between the Great Blue Formation on the south structural block and the Deseret Limestone. While no fault exposures are known, some breccia zones several feet thick are present and also the omission of more than 700 feet of the Humbug Formation. In addition, only a partial thickness of the other two formations, and the contact relationships suggest a buried fault between the two formations. The thrust fault is likely the result of west to east compression related to the development of the Lake Mountain Syncline during the Sevier compressional orogeny.

west

A longitudinal fault occurs along the axis of the subsidiary anticline. The main length of the fault is in the quadrangle to the north. Within the map area, however, the western wall of the fault consists of Humbig Formation and the eastern hanging wall of Great Blue Formation. Displacement is normal, and the fault plane is vertical to steeply east dipping with the east wall downed way. Veins of calcite and aragonite to more than eight feet thick occur along this fault zone, and are the basis and underground calcite mine in the area. The north trending normal fault is not genetically related to the fold, the tear fault or the thrust fault. The normal fault appears to have developed synchronous with other Basin and Range faults at a later time.

Okerlund (1951) suggests that the fault where hosts the calcite vein is likely associated with Basin and Range block faulting. Bullock (1949) describes the Lake Mountains as a fault block range with an east-bounding fault near the shore of Utah Lake. The fault noted above is less than a mile from the lake shore and tends to parallel it. Evidence for block faulting is suggested by the relative position of the fault to Utah Lake and the uplifted mountain block to the west.

In section 6, in the Little Cove area, sec 6, T7S, R1E, a gravity slide fault of north trend occurs in the partly cemented Lake Bonneville gravels. They dip about 35 degrees west and strike north. The deposits appear to have been rotated westward on a listric-type fault of moderate to steep east dip. This is latest structural dislocation noted in the area, and is probably less than 15,000 bp. the fore present by

Lincoln Point

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The Lincoln Point block (LPB) consists of strata of the Oquirrh Formation and is a structural segment of the large overturned West Mountain anticline Bissell, 1948; White, 1953). The eastern part of the structure is an overturned limb which dips westward at about 50 degrees. To the South and West the beds are upright and dip eastward about 15-30 degrees. A north-trending westward inclined reverse fault may lie? between the west and eastern parts of the structure. If present at is covered by Lake Bonnevelle deposits in the valley area. Such a fault is inferred based on the lithologic differences in the units exposed in the area. The northern extension of West Mountain is a horst block as the based on the uplifted character of the mountain block, the abrupt termination of bedrock on either side of the block, and known faults, and associated thermal springs along both sides of the mountain block.

The overturned anticline plunges NNEward under Utah Lake sediments and probably continues northward under Bird Island. The Nature consists of Holdcene tufa. Cottam (1926) suggests a guartzite pedestal is the foundation for the tufa deposits of Bird Island. Brimhall (1979) Suggests that the Bird Island Fault lies along the west side of Lincoln Point and Bird Island, and that an unnamed fault occurs east of the island and just east of Lincoln Point. He also suggests a down to the west displacement on the Bird Island Fault and down to the east movement on the unnamed eastern fault based on sonic measurements. Based on these apparent displacements and the relatively recent fault displacements in Utah Lake sediments east of Lincoln Point, Brimhall (1979) suggests orcludes that the fault block of northern West Mountain and Bird Island stands structurally high as a horst block. Known tufa deposition and thermal spring activity are probably related to these faults.

# Lake Unit

In the regional setting Utah Lake is a fault basin lying between mountain masses of Lake Mountain, West mountain, the Wasatch Mountains and the Traverse Mountains. Each of these mountain masses is considered to be a fault block range related to Basin and Range faulting produced under extensional stresses, ( ). The Lake unit area covers a large part of the Lincoln Point Quadrangle. Brimhall (, 1979 ) thoroughly investigated this part of Utah Lake and outfined several significant structural feactures present in the relatively recent lake sediments. We summarize his form

0 Step !! Brimhall (1979) and his assistants, identified a series of normal faults of north to northeast trend within the lake sediments, of the Lake unit. The normal faults show as small step faults ( Fig. , 1, 3 ); a graben (2) and a horst (4) along the northwest-southeast profile. The horst is a north extension of the Lincoln Point block under the lake. Characterestics of the faults are discussed separately below Bird Island Fault Within the quadrangle the Bird Island Fault extends northeastward from the west side of Lincoln Point to the west side of Bird Island and then north northeastward off the map area. The fault is at least ten miles in length. The acoustical Source profile of Brimhall ( 1979 ) indicates the western wall of the fault is downthrown relative to the eastern wall and the displacement of the Holocene deposits is as much as two meters ( 6.6 ft ).

Either an eastern branch or intersection of another fault occurs on the Bird Island fault about three kilometers (1.9 miles) north of Bird Island. This eastern fault extends southward and east of Bird Island toward the east side of Lincoln Point. It is possible that the fault is part of a fault zone along the east side of West Mountain and may represent relatively recent movement in this zone. The eastern block is downthrown about two meters relative to the western block. The block between the Bird Island Fault and the east fork fault thus is a structural horst.

East Goshen Bay Fault --- The southern part of the East Goshen Bay Fault (EGB) is parallel to and a few handred feet west of the Bird Island Fault. The EGB fault extends northward for about 10 miles to the quadrangle boundary. Near the center of the map a branch of the fault trends northerly to the map edge. South of the branch intersection the western block is down relative to the eastern block. North of the intersection the block between two faults is downthrown or graben-like in charater, Brimhall (1979) named this the Pelican Point Graben. It represents the lowest structure of Utah Valley. The acoustical profiling (Fig. ) shows a displacements of at least 5 meters, (16 Feet ) on the EGB fault. The fault cuts the clay unit of the Provo Formation of uppermost Pleistocene, and is the age of the last displacement on the fault.

West Goshen Bay Fault --- The West Goshen Bay Fault (WGB) has a strike length of about 6 miles (9.6 km) and parallels the northwest shore of Utah Lake in the map area. It extends southwestward and northeastward into adjacent quadrangles. The eastern block of the fault is downthrown as is the fault displacement on the EGB fault and thus forms the Goshen Bay Graben (Fig. ).

Each of the above foults is signif, can't on that they indicate the recency of feult moment in the lefon valley and the displacement in feet (m).

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Ground Water

Price and Conway (1988) report on the ground water resources of the southern Wasatch Front area, of which Lincoln Point quadrangle is a part. Over a 20 year period, 1966-1986, the altitude of the potentiometric surface of the flowing wells rose 5 feet. However, in the basin fill near Benjamin a decline of 4 feet in the water level occurred in the same period. Apparently little change in the water level occurred to the east of Lincoln Point in the valley fill.

Cordova (1970, p. 56-57) estimates that the volume of theoretically recoverable water in the upper 400 feet of saturated basin fill in southern Utah Valley and Goshen Valley totals about 6 million acre feet. In terms of transmissivity of the basin fill the southeast Lincoln Point valley fill lies in the range of 100-10,000 feet square per day (i.e., cubic feet per day per foot). This is shown on a map by Price and Conroy (1988, Plate I). The authors (Plate 3) also suggest the dissolved solids concentration in waters in the basinfill in the same area is in the low range of 100-500 mg/l, in strong contrast to the 1000-3000 mg/l on the westside of West Mountain from Lincoln Point south to Goshen. In the Lake Mountains, part of the map area, two wells average 1525 mg/l dissolved solids.

Thermal Springs-Three thermal springs flow in the map area. Two are correct south of Goose Point in the northwestern part of the area near the west shore of Utah Lake and about a mile apart. The third is spring is on Lincoln Point on the lake shore in section\_\_\_T R\_\_\_.Another thermal spring occurs just north of the map boundary in section \_\_T R\_\_. The spring flow is small, and temperatures approximate 98-100 degrees fahrenheit.

Several drilled wells also contain warm water. Figure \_\_shows the position of \_\_warm water wells just north of the map border. Temperatures range from \_\_ to \_\_\_.

Composition of the warm water wells compared to fresh water wells is shown in Table \_\_\_\_. The fresh water (Hem, 1985) contains less than 1000 mg/l of dissolved solids, and 1000-3000 mg/l is considered slightly saline, while <10,000 mg/l is moderately saline, briney is < 35000 mg/l of dissolved solids. All samples collected within the quadrangle are slightly saline or less.

Surface Water-Utah Lake has a usable storage capacity (Price and Jensen, 1982) at the highest legal storage level of 883,900 acre feet (an acre foot is 1 foot ofwater over 1 acre). The lake receives an average of 602,000 acre feet per year from the major in-flowing streams such as the Provo River 303,000, Spanish Fork 155,000. Gaged Wasatch Front streams 28,300 acre feet and ungaged streams 15,200 acre feet (Hyatt and others, 1969). In the lake springs supply additional water. The authors conclude that the surface waters

The authors conclude that the surface waters range from fresh to slightly saline. Locally some ponds and lakes on the southwest side of West Mountain may exceed 1000 mg/1, and also Utah Lake at specific times. Spanish Fork River may also exceed 1000 mg/1 in its lower reaches during runoff time. This river enters Utah lake about 4 miles N70E of Lincoln Point. Benjamin Slough to the south shows a range of 500 to 1000 mg/1 in runoff waters based on

specific conductance reading at the measuring site slightly south of the map boundary in section 14, R1E, T8S.

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